Patterns of Microbial Heterotrophy Through Changing Environments in a Marine Sediment

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ABSTRACT: Heterotrophic activity was measured through a vertical profile in Halifax Harbor sediments using 14 C labeled glucose, glutamate, and lactate as substrates. A modified heterotrophic potential method was used to detect the amount of label that was converted to low molecular weight metabolites as well as the amounts respired to CO_2 and incorporated into cellular material. For all three substrates, total uptake was greatest at the transition from aerobic to anerobic sediments (40 cm subbottom). The percentage of the total uptake respired reflected the predominant type of heterotrophic metabolism at the various horizons with the highest values occurring above (aerobic respiration) and below (sulfate reduction) the 40 cm horizon. Incorporation of the substrates into cellular material was low for both glutamate and lactate (< 5 %) but averaged approximately 30 % for glucose. For all three substrates tested, low molecular weight metabolites represented a significant portion of the substrate uptake at many horizons (approximately 50 % at 40 cm). Data are presented indicating that the observed increase in heterotrophic activity at 40 cm sub-bottom may be due to chemoautotrophic organic carbon production at this horizon.

INTRODUCTION

Comparatively little is known about the biological processes in sediments even though they account for a major portion of coastal ecosystems. With bacterial populations as high as 10^{10} g⁻¹, the sediment microbiota outnumber, and may be as important as, the macrobiota. Despite this, microbiology is often the least understood and studied discipline of marine science. Research in sediment microbiology has been focused on the determination of biomass, the nitrogen cycle, sulfate reduction, and methanogenesis. For examples see Erkenbrecher and Stevenson (1975, 1978) and Chrzanowski et al. (1979); Koike and Hattori (1978a, 1978b), Sørensen (1978); and Oren and Blackburn (1979); Cappenberg (1974a, 1974b), and Winfrey and Zeikus (1977), respectively. Only a limited number of studies have measured heterotrophic activity per se (Hanson and Gardner, 1978; Meyer-Reil, 1978; Meyer-Reil et al., 1978). In addition, there has been only one report on the type and amount of microbial heterotrophic activity through a vertical sequence of marine sediments, where Sørensen et al. (1979) have measured aerobic and anaerobic respiration in the upper 10 cm of a coastal marine sediment by following oxygen concentrations, nitrogen speciation, and 35SO₄

reduction. Moreover, while studies have been reported that measure the aerobic (Meyer-Reil, 1978; Meyer-Reil et al., 1978) and anaerobic (Hanson and Gardner, 1978) uptake of organic sustrates, there has been no description of heterotrophic activity through a gradient of oxygen concentration such as one is apt to find in most marine sediments.

Recently, Kepkay et al. (1979) and Kepkay and Novitsky (1980) have described autotrophy as well as some physical and chemical parameters in a coastal marine sediment. These sediments are characterized by a relatively deep aerobic zone, a well defined aerobic to anaerobic transition, and an anaerobic zone still within easy reach of sampling. The purpose of the present study, therefore, was to characterize the heterotrophic microbial activity of these sediments under simulated *in situ* oxygen concentrations using a radioisotope uptake technique that was modified to monitor aerobic as well as anaerobic activity.

MATERIALS AND METHODS

Sampling

Diver-collected cores were obtained from Halifax Harbor (Canada) as previously described (Kepkay et al., 1979; Kepkay and Novitsky, 1980) and stored on ice during transport. In the laboratory the cores were split

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and subsamples were taken from the interior of the core at various horizons with modified 5-ml plastic disposable syringes. The subsamples were immediately placed in serum bottles and flushed with mixtures of oxygen and nitrogen. Kepkay and Novitsky's procedures were followed in preparation of the gas mixtures.

Measurement of Heterotrophic Activity

Relative heterotrophic activity was measured by the single concentration of substrate technique of Griffiths et al. (1977) using ¹⁴C-(UL)-glucose, specific activity 360 mCi mmol⁻¹; ¹⁴C-(UL)-glutamic acid, specific activity 296 mCl mmol-1; and 14C-(UL)-lactic acid, specific activity 138.6 mCi mmol⁻¹. Sediment samples (1-2 g) were diluted 1:100 in filter sterilized bottom water and 10-ml portions of this mixture were placed into serum bottles. Labelled substrate was added to each (0.05 μ Ci) and the bottles were again capped and flushed with the appropriate air-nitrogen mixture to simulate the in situ O_2 concentrations of the sediments. In all cases the amount of labelled substrate added to each bottle exceeded the unlabelled substrate in the slurry. Incubation was at the in situ temperature of 2 °C for 2 h before the addition of 1 ml of 3 N H₂SO₄ to stop the reaction. The contents of each incubation were then filtered through Millipore 0.45 µm filters to collect high molecular weight material. The low molecular weight organics in the filtrates were stored at room temperature and pH 1.0.

Determination of Low Molecular Weight Metabolites

The amount of substrate that was converted to low molecular weight metabolites was determined by thin layer chromatography of the filtrates from the heterotrophic activity determinations. In general, 100 μ l of the filtrate was spotted on a thin layer plate next to the spot of the control filtrate (acid killed at time O). After running the chromatogram in the appropriate solvent, the control lane was scraped off in 0.5-cm bands which were placed into scintillation vials and assayed for radioactivity. Once the peak of residual substrate was located, the corresponding area of the test chromatogram was scraped into a scintillation vial. The remaining chromatogram was scraped into another vial and both were assayed for radioactivity. Another control was treated in like manner to determine the background of non-substrate label present. After subtraction of this blank, the ratio of residual substrate to the modified substrate was calculated and the amount of low molecular weight metabolites generated in the course of the incubation was then determined.

Glucose filtrates were run on cellulose coated plates with a solvent consisting of ethyl acetate-pyridinewater, 40:20:40, upper phase only.

Glutamate filtrates were run on silica gel coated plates with a solvent consisting of n-butanol-acetic acid-water, 80:20:20.

Lactate filtrates were run on silica gel coated plates with a solvent consisting of ethanol-chloroform-ammonium hydroxide-water, 53:303:15.2:1.5

All chromatograms were run at room temperature and the solvents were allowed to travel 10 cm.

Calculations

The total amount of substrate taken up was calculated as the sum of the amount of label in cellular material, CO_2 , and low molecular weight metabolites. Individual components were then expressed as a percentage of this total.

Methane Determination

Pore water was extracted from the sediments as previously described (Kepkay and Novitsky, 1980). Dissolved methane was measured relative to a seawater standard saturated with methane at O °C and 1 ATM. The procedure used was the direct injection of 25 μ l of pore water or standard water onto a molecular sieve (MS 5A) column in a gas chromatograph with an electron capture detector (described by Cooke, 1973). The maximum standard deviation of the methane determinations was 0.002 μ moles CH₄ ml⁻¹.

RESULTS

The relative uptake profiles for the 3 substrates are shown in Figures 1a, 2a and 3a. Data are presented for one representative core only. Data were collected for 2 other cores collected in the same area but on different dates. Comparable uptake profiles from different cores were similar, all showing maximum uptake at the aerobic to anaerobic transition. For all three substrates, the uptake profile is characterized by a single peak of elevated activity at 40 cm sub-bottom. The glucose profile was assayed at finer depth intervals and showed the same general pattern as glutamate and lactate. Intermediate peaks were not apparent at these finer intervals. The data are presented in Figures 1 to 3 with and without the inclusion of the low molecular weight metabolites produced during the incubation. As can be seen, the inclusion of this material significantly changes the magnitude but not the general

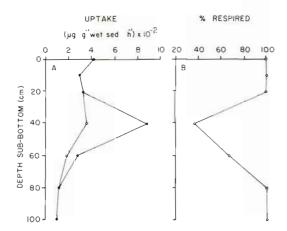


Fig. 1. Uptake and respiration of ¹⁴C lactate by microbial populations through a vertical profile of Halifax Harbor sediment. Uptake (A) was calculated in the usual manner () and with the inclusion of low molecular weight metabolites produced during the incubations as determined by chromatography (•). The percent respired (B) was calculated with the inclusion of the low molecular weight metabolites

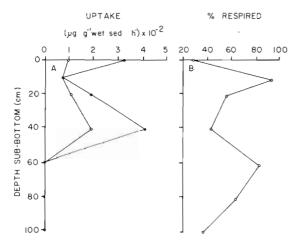


Fig. 2. Uptake and respiration of ¹⁴C glutamate by microbial populations through a vertical profile of Halifax Harbor sediment. Uptake (A) was calculated in the usual manner (·) and with the inclusion of low molecular weight metabolites as determined by chromatography (•). The percent respired (B) was calculated with the inclusion of the low molecular weight metabolites

character of the uptake profile. The changes are most pronounced with lactate uptake at the 40 cm horizon (Fig. 1a) and with glutamate uptake at the 0 and 40 cm horizons (Fig. 2a). The uptake of both glucose and glutamate show a marked decrease (i. e., of up to 2 orders of magnitude) below 40 cm. When sediment slurries from various depths were incubated under the high oxygen concentrations of room air, the resultant uptake profiles were similar in shape but several times greater in magnitude than those produced under simulated *in situ* oxygen conditions.

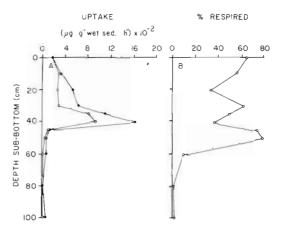


Fig. 3. Uptake and respiration of ¹⁴C glucose by microbial populations through a vertical profile of Halifax Harbor sediment. Uptake (A) was calculated in the usual manner (·) and with the inclusion of low molecular weight metabolites as determined by chromatography (•). The percent respired (B) was calculated with the inclusion of the low molecular weight metabolites

The percentage of the substrate taken up that was converted to cellular materials was low for both lactate and glutamate (less than 5 %), but averaged about 30 % in the upper horizons when glucose was the substrate (Fig. 3b). In contrast, the percentage of the substrate taken up that was respired to CO₂ was very high for lactate (Fig. 1b) and glutamate (Fig. 2b) and moderately high for glucose (Fig. 3b). The percentage of lactate respiration shows a dramatic drop at 40 cm with nearly 100 % of the label respired above and below. With glutamate as the substrate there was also a drop in the percent respired at 40 cm. However, the results for glucose respiration do not show such a pronounced effect. For both glutamate and glucose, the values for respiration below 40 cm are shown, but with the total uptake very low it is questionable whether these figures are valid. The amount of substrate converted to low molecular weight metabolites varies from 0 to 70 % for all three substrates. Since the incorporation of labelled lactate or glutamate into cellular material was low, the results for conversion to low molecular weight metabolites are essentially a mirror image of the respiration results. When lactate was used as a substrate, low molecular weight metabolites were only detected at 40 and 60 cm, while for glutamate these metabolites were found throughout the profile and were especially high at 0 (70 %) and 40 cm (54 %). With glucose as the substrate, low molecular weight metabolites varied from 0 to 50 % with no apparent pattern throughout the profile. As an example, the breakdown of substrate utilization at the 40 cm horizon for all 3 substrates is shown in Table 1.

A profile of dissolved methane in the pore water of these sediments is shown in Figure 4. No methane was

Table 1	Fate of	labelled	substra	ate taken	up b	by the microbial
pop	ulation	from the	40-cm	transition	ıal e	nviroment

Substrate	Percentage of substrate taken up associated with:					
	Respiration	Cellular macro- molecules	Low molecular weight metabolites			
Glucose	36.4	21.0	42.6			
Glutamate	43.6	3.3	53.1			
Lactate	36.3	4.2	59.5			

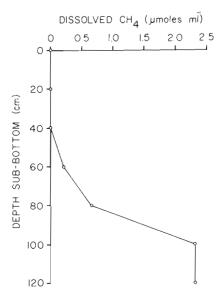


Fig. 4. Profile of the concentration of dissolved methane in pore water of Halifax Harbor sediments. Methane concentrations were determined by gas chromatography. Twice the maximum standard deviation (0.002 μ moles ml⁻¹) is smaller than the diameter of data points

detected above 40 cm; below 40 cm methane was detected in increasing amounts down to 100 cm where it reached saturation levels.

DISCUSSION

To assess the relative microbial activity in these sediments we chose the single substrate concentration method of Griffiths et al. (1977). This procedure has well-defined limitations. The concentration of labelled substrate is usually in excess of the unlabelled substrate concentration so that the potential rather than the actual uptake rate is determined. However, the method can be used to compare heterotrophic activity in profile through a sediment. With this in mind, the V max of substrate uptake calculated by Kepkay and Novitsky (1980) and Novitsky and Kepkay (unpub-

lished), using Griffiths et al.'s (1974) modification of the technique of Wright and Hobbie (1966), compare well to the data presented here. An attempt was made to determine the uptake rate using undisturbed sediment cores according to the method of Meyer-Reil (1978), but the compact nature of these fine-grained sediments prevented the infiltration of substrate even when high vacuum was applied.

Previously, all methods for the determination of the heterotrophic activity of seawater or marine sediments ignored the possibility of conversion of labelled substrate to anything but CO₂ or cellular material. This may be the case in many aerobic systems. However, it seems likely that in areas of reduced oxygen concentration or in anaerobic areas that fermentation end products would constitute a large portion of the substrate uptake as has been demonstrated by Christian and Wiebe (1978) for salt marsh soils. Using the chromatographic method described here, we were able to detect large amounts of labelled material that were not retained by filtration, i. e. low molecular weight metabolites. Due to the chemical heterogeneity of this material we were unable to identify its components. It should contain both intracellular (pool material and low molecular weight metabolites) and extracellular (metabolic and products and leakage) material. Measurements of this kind are of obvious importance because this material would contribute directly to the pool of dissolved organic carbon in the natural environment.

The gross uptake of the three substrates by these sediments is characterized by a large peak of activity at the transition from aerobic to anaerobic sediments (40 cm sub-bottom). This is also the area where a peak in bacterial cell numbers (determined by acridine orange direct counts) occurs (Kepkay et al., 1979). The specific activity index proposed by Wright (1978) attempts to correct activity values for changes in population size by dividing the measured activity by bacterial numbers. The resulting index should give an estimate of activity based on factors other than population size. We have calculated the indices for glutamate, glucose and lactate uptake and have found that only the uptake of lactate is closely related to population size. The specific activity of lactate uptake is independent of depth, indicating uniform uptake per unit population throughout the sediments examined. However, bacterial numbers in themselves are often no indication of a population's activity. Kepkay et al. (1979) and Kepkay and Novitsky (1980) have shown that this transitional environment is also active with respect to the chemoautotrophic production of organic carbon. Using our previous methods (Kepkay et al., 1979; Kepkay and Novitsky, 1980) for detecting CO₂ fixation, but stopping the reaction with rapid filtration rather than acidification, it was found that 49 % of the organic carbon fixed by populations from the 40 cm horizon was present extracellularly (Novitsky and Scott, unpublished). This means that there is a constant production of dissolved organic matter available for microbial populations. We feel that this is the major, if not the paramount, factor responsible for the elevated cell numbers and activity at this horizon. It must be emphasized, however, that the contribution by bacterial biomass to the organic carbon content of the sediment at this and other horizons is very small. The bacterial biomass contributes little (i. e., less than 1 %) to the organic pool (Kepkay and Novitsky, 1980) and as such it cannot greatly influence the biology and/or chemistry of these sediments. This small biomass does not preclude a significant bacterial role in biochemical transformations. As an example, we have shown the chemoautotrophic generation of organic carbon at the 40 cm horizon to be 1.4 imes 10⁻² μg g^{-1} wet sediment h^{-1} (Kepkay and Novitsky, 1980). When the sedimentation rate of only 2.5 mm yr⁻¹ is considered, the residual organic carbon at 40 cm is a reflection of long-term accumulation in situ rather than the expression of a point-in-time population of bacteria.

Below the 40-cm horizon, uptake activity under simulated in situ anaerobic conditions for glucose and glutamate falls by 2 orders of magnitude indicating that these are probably not the normally encountered substrates in vivo. Stimulation of glucose and glutamate uptake by the addition of oxygen indicates that populations are present which are capable of taking up these substrates even though they may be dormant under in situ conditions. The horizons immediately below 40 cm are active with respect to sulfur reduction (Kepkay and Novitsky, 1980) and exhibit active uptake of lactate, a preferred substrate for sulfur reducing bacteria. Lactate, other fermentation end products, and short chain fatty acids can all be converted, through a number of steps, to methane (Cappenberg, 1976). The eventual production of methane may also be contributing to the sustained uptake of lactate below 60 cm. Since strict anaerobic techniques (Hungate, 1969) were not followed, the only evidence for methanogenesis is the presence of methane (Fig. 4). Again, this may be more indicative of the long-term accumulation of end products rather than intensive short-term production.

The only apparant anomaly in uptake patterns is the enhanced uptake of glutamate just below the mudwater interface. This enhanced activity can only be detected then the low molecular weight fraction is included in the uptake calculations. Quite possibly, the cells – limited with respect to nitrogen as suggested by Kepkay and Novitsky (1980) – are deaminating the glutamate for a nitrogen source and discarding

the carbon skeleton. Because there are many other possibilities, the exact reason for this enhanced uptake remains unknown.

The data reflecting the fate of the substrate after uptake demonstrate that a pattern of microbial metabolism exists that is related to the various environments encountered in vertical section through these sediments. Incorporation of lactate and glutamate into cellular carbon was very low (< 5 %). This is especially surprising in the case of glutamate, which is an amino acid that one would expect to be easily incorporated into protein. Incorporation of glucose is much higher (averaging 30 %) and may be due to the versatility of this central metabolite. The presence of soluble low molecular weight material was most pronounced in and near the aerobic-anaerobic transition and, assuming the amount of intracellular pool and metabolite material to be small, most of this carbon would therefore be excreted metabolic and products. Considering the reduced amount of oxygen at the transition, one would expect fermentation to be the predominant type of metabolism in this area. This conclusion is supported by the fact that the percent respired for all three substrates is low at the horizon. Above and below the transition, the percent respired for glutamate and especially lactate is nearly 100 %,

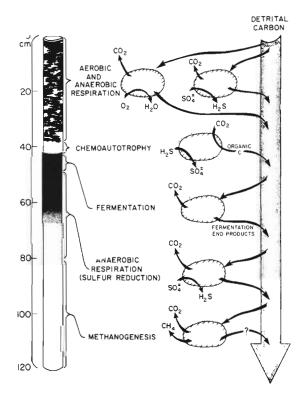


Fig. 5. Schematic diagram of microbial metabolism and carbon flow in Halifax Harbor sediments. Only the predominant type of metabolism is listed for the various horizons. This does not imply that they are exclusive or nonoverlapping

indicating aerobic respiration and anaerobic respiration (sulfate reduction), respectively. The percent of glucose respired follows a somewhat similar pattern with a low value at 40 cm and higher values immediately below; this indicates that CO2 is probably produced at the aerobic to anaerobic transition by predominately fermentative metabolism. A clear explanation of the percentage of glucose respired in the overlying aerobic zones is masked by the relatively large amount of substrate that is converted into cellular material. Our data do not compare well with those of Meyer-Reil (1978) (an average of 8 % of aerobic glucose uptake respired) or Hanson and Gardner (1978) (20 to 50 % of anaerobic alanine uptake respired and 5 to 20 % of anaerobic asparate uptake respired). The reasons for the large discrepancy probably reflect the diverse types of sediments studied: a sandy beach (Meyer-Reil, 1978), a salt marsh (Hanson and Gardner, 1978), and the ponded muds of this study. It is clear then that sedimentary environments must be taken into account before any meaningful comparison of microbial activity in these environments can be made.

The general succession of microbial metabolism in the Halifax Harbor sediments is summarized diagrammatically in Figure 5. One of the primary factors governing this succession is the availability of oxygen and sulfate as terminal electron acceptors. The various groups of heterotrophs not only mineralize the organic carbon pool but are also active in changing its composition through their metabilism. The constant production of organic carbon by chemoautotrophy at the transition between aerobic and anaerobic sediments fuels the active heterotrophic populations. Actual bacterial biomass, as evident in direct measurements and the incorporation of labelled substrates, is insignificant with respect to the organic carbon pool in these sediments.

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